

Workshop Hydrological Optics :



Measurements and Modelling

Convidado Externo – Emmanuel Boss – University of Maine

Projeto: FAPESP 2015/19653-0

Organizado pelo

Laboratório de Instrumentação de Sistemas Aquáticos (LabISA) – INPE/OBT

19 e 20 de abril de 2016 10 às 18:00 horas

Auditório do Instituto Interamericano para Pesquisas em Mudanças Globais - IAI Auditório José Simeão de Medeiros – LabGeo

Sala 27 - Prédio Asa

Instituto Nacional de Pesquisas Espaciais - INPE

São José dos Campos - SP

Participantes

















In situ radiometric and bio-optical measurements at the Lower Amazon River

Aline de Matos Valerio Milton Kampel

19-20, April 2016

TROCAS



 This work is part of the Net Ecosystem Exchange of the Lower Amazon project (TROCAS), that seeks to unravel the sequences of processes and sources of terrestrially-derived organic matter that culminate in the immense CO₂ outgassing to the atmosphere.

- IOP and AOP
- AOP:
 - Remote Sensing of Reflectance (R_{rs})
 - Diffuse attenuation coefficient (Kd)
- IOP:
 - Absorption Coefficient
- Optical closure and the develop of an algorithm for the color of the water
- Lower Amazon River
- Preliminary results
- Discussion and ideas

IOP and AOP

 The color of the water is determined by the scattering and absorption of incident light by the optically active constituents (OAC) as well as by the very pure water





Remote Sensing of Reflectance

 The Remote Sensing Reflectance (R_{rs}) it takes into account the reflected sky radiance as well the wind speed and the viewing geometry.



Remote Sensing of Reflectance

- Residual glint correction
- Mobley (1999) and Ruddick (2006) suggest to do the residual glint correction.
- Garaba et al. (2013) discuss that there is no consensus of the best residual glint correction.
- Is it possible (or right) to correct in optically complex waters with highly scattering optically active materials?

K_d and Secchi Disk

Secchi depht is approximately inversely proportional to the vertical attenuation coefficient for downwelling irradiance (K_d)

Koenings and Edmundson (1991)

- In inland waters the relation between K_d and Secchi Disk is not constant. Secchi depht is particularly sensitive to turbidity.
- For a given sun angle K_d can be considered as a IOP or quasi-IOP Kirk (2011)

$$K_{d}(z, \lambda) = -\frac{1}{E_{d}(\lambda)} \frac{dE_{d}}{dz}$$

Mobley (1994)

Bio-optical properties

- Absorption coeficiente of:
 - Chromophoric Dissolved Organic Matter (CDOM)
 - Spectral parameter: Slope of de CDOM
 - Suspended particulate matter
- NAP Phytoplankton
- Chlorophyll-a concentration (chl-a)



Chromophoric Dissolved Organic Matter





Filter 0.2 µm

Chromophoric Dissolved Organic Matter

- Slope of the CDOM
 - The absorption of the CDOM is typically modeled as an exponential function in which a spectral slope parameter (S) describes de rate of decrease in absorption with increase of wavelength.



Nelson & Siegel, 2013

Suspended particulate matter

A collective name for solid particles in the water which is trapped on a 0.7 μ m filter.



ap(total)

Chlorophyll-a concentration



•Chlorophyll is a key biochemical component that is responsible for photosynthesis.

•Chlorophyll is essential to the existence of phytoplankton. Phytoplankton can be used as an indicator organism for the health of a particular body of water.



Filter 0.7 µm

- There are several studies that estimate and monitor the bio-optical properties by remote sensing in waters dominated by phytoplankton, as the oceans. However, the use of remote sensing to study the bio-optical properties of optically complex waters such as rivers, faces some challenges that must be overcome.
- From an optical point of view, the lack of a direct relationship between the NAP absorption and the phytoplankton makes difficult the general parameterization of the optical properties. Thus, bio-optical algorithms develop for ocean applications can peform poorly when applied in inland waters.
- Thus, it is necessary to develop or adjust algorithms proper to inland waters and estimate the COA individually. The difficult of estimating individually COA comes from the similarity of the spectral curve of NAP and CDOM absorption. Therefore, new algorithms are based on diferences in their backscattering curves in the visible region near 555 nm (Matsuoka et al., 2013; Zhu, Yu, 2013) or the relationship between the ratio of Kd and aCDOM (Loisel et al., 2014).

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• Empirical (band ratios)

Band ratio Regression Less knowledge about IOP X AOP Limited to similar Waters

Semi-analytical

Matrix inversion methods

Empirical (band ratios)

Semi-analytical

RTE IOP X AOP Different types of water Needs empirical relationships

Matrix inversion methods

Empirical (band ratios)

Semi-analytical

Matrix inversion methods

Uses semi-analytical method Require knowledge of specific IOP (SIOPs) Generally not applicable across different environment without field measured SIOPs

Presentation Objective

- Present the in situ radiometric and bio-optical data collected until the moment to subsidize the discussion about my PhD goal: develop na algorithm to estimate CDOM and NAP separately.
- Why estimate CDOM and NAP separately?

Study Area



Despite of the Lower Amazon region be hydrologically relevant as it has already received the contribution of all major tributaries of the Amazon River, there is still a gap in the remote sensing literature.

Study Area



Study Area



Mean discharge (2005-2015)

Discharge data obtained at ANA station in Obidos



- Radiometric measurements:
 - Above water: FieldSpec ASD
 - In water: Satlantic Profile
- Residual Glint Correction:
 - Kutser et al. 2013
- Ratio of NIR wavelenghts:
- Turbid waters (Ruddick et al., 2006)
- Secchi Disk



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Fluorimeter for chl-a Backscattering 470 and 700 nm Temperature

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Problems: Highly turbid water Fluorimeter doesn't work - too much sediments Bb (470 nm and 700 nm) doesn't work: saturates Extinction of the light in the first meters (~3 m) Too much tilt in the first meters.



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🔶 Secchi Disk

Problem: Strong waves at Macapa and too turbid water



- Absorption Coefficients:
- The filtered samples were storaged at ultrafreezer during the cruise.
- After the cruise the samples of absorptions coefficients were read using a spectrophotometer.
- A bench fluorimeter was used to read the chl-a.

















Т2

Т3

T1













Problem: Besides strong waves at Macapa and too turbid water, all the Lower Amazon region are influenced by the tide and sometimes the Amazon River intrudes others rivers in certain times of the day. Uncertainty due the visual inspection

• Satlantic data: Still working on!

K_d(PAR)



 $K_d(PAR) > 2 m^{-1} = turbid waters$

Kirk (2011) and Koenings and Edmundson (1991)





Max: 4.5

Min : 0.5









chl-a(mg m⁻³)

High: 10			
		Min	Max
	T1	0.39	28.29
	T2	1.25	8.80
	T3	0.39	3.78

Low: 1

























Chl-*a* concentration (mg m⁻³) Min 0.39 Max 3.78



TROCAS 3

aCDOM aphy aNAP 2,5 2,0 1,5 T1 1,0 0,5 0,0 400 450 500 550 600 650 700 750 800 1,5 T2 0,5 400 450 500 550 600 650 700 750 800 T3 400 450 500 550 600 650 700 750 800



Other measurements

- Conductivity
- DO
- Temperature
- pH
- TSS
- DOC
- DOM
- POM

- pCO2
- DIC
- Methane
- Nutrients
- ADCP
- Etc.....

- The Rrs were all very characteristic of highly turbid waters, and the Rrs of Tapajos and Xingu it has phytoplankton signature.
- Preliminary analysis of this data set shows that during the high water stage, the Amazon River is dominated by CDOM and detritus, while the Tapajós and Xingú rivers are dominated by chl-a and CDOM.
- During the falling and low water stage, the absorption coefficients are well mixed at all the stations, with a slightly high chl-*a* at the Tapajós River.
- It was possible to see the HIGH VARIABILITY of the bio-optical properties in waters of the Lower Amazon River during three discharge water stages: high, low and falling water.

- Correction of residual glint?
- Semi-analytic algorithm?
- Matrix inversion method?
- Empirical algorithm?
- Other?

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TROCAS TEAM and collaborators

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